New Particle Growth and Shrinkage Observed in Subtropical Environments

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New particle growth and shrinkage observed in subtropical environments

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Abstract. We present the first systematic analysis for new particle formation (NPF), growth and shrinkage of new particles at four different sites in subtropical central Taiwan. A total of 14 NPF events were identified from 137 days of ambient measurements during a cold and warm season. The measured formation rates of 10 nm particles \( J_{10} \) and growth rates were in the range of 4.4–30 cm\(^{-3}\) s\(^{-1}\) and 7.4–24 nm h\(^{-1}\), respectively. The onset of NPF events coincided with decreases of condensation sink (CS) and increases of SO\(_2\) under enhanced atmospheric mixing and dilution. However, the lower or comparable SO\(_2\) on event days than on non-event days suggests that SO\(_2\) was not a limiting factor for NPF. On non-event days, the particle number concentrations were mostly driven by traffic emissions. We also observed shrinkage of new particles, the reversal of growth, during five out of the identified secondary formation. UFP particles events. In intense cases, the grown particles shrank back to the smallest measurable size of \( \sim 10 \) nm, thereby creating a unique “arch-like” shape in the size distribution contour plot. The particle shrinkage rates ranged from \(-5.1\) to \(-7.6 \) nm h\(^{-1}\). The corresponding particle volume losses suggest that a notable fraction of the condensable species that contributed to growth was semi-volatile. The particle shrinkage was related to enhanced atmospheric dilution, high ambient temperature and low relative humidity, thus favoring the evaporation of semi-volatile species from the particulate phase to the gas phase. Our observations show that the new particle growth could be a reversible process, in which the evaporating semi-volatile species are important for the growth of new particles to sizes of environmental health concerns.

1 Introduction

Atmospheric aerosols impact various aspects of human health, air quality and climate. Aerosol particles are also linked to short- and long-term cardiovascular disease and mortality (Atkinson et al., 2010; Pope et al., 2009). Fine particles of diameters smaller than 2.5 micrometer can be more pertinent to health effects (Peng et al., 2009; Zanobetti and Schwartz, 2009). And, ultrafine particles (UFPs) in the size range of smaller than 100 nm can be even more toxic, as they deposit more efficiently in the alveolar region than larger...
particles (Nel, 2006). In ambient air, aerosol number concentrations are dominated by these UFPs.

Atmospheric UFPs are from two source processes, primary emissions and secondary formation. Primary UFP particles are directly emitted from motor vehicles, particularly heavy-duty diesel engines (Morawska et al., 2008). For example, emission studies of heavy-duty diesel engines have shown that non-volatile UFPs emitted from the exhaust are on the order of $10^{7}$–$10^{8}$ cm$^{-3}$ (Heikila et al., 2009; Young et al., 2012), several orders of magnitude higher than ambient UFP concentrations. The secondary formation process that governs particle number concentrations is new particle formation (NPF). NPF involves nucleation (gas to particle conversion) and initial growth to sizes of 1–3 nm detectable by current instruments (Kulmala et al., 2012; Vanhanen et al., 2011). Once clusters formed via nucleation and become larger than the critical size, growth becomes spontaneous. Subsequent growth of freshly nucleated particles often persists for many hours under favorable atmospheric conditions (e.g., under low pre-existing aerosol particle and high concentrations of low volatility compounds) to become larger than ~50 nm and could act as cloud condensation nuclei (Dusek et al., 2006). Unlike motor vehicle emissions that are limited to urban areas, NPF events have been observed in all types of environments, including rural, coastal, urban areas and boreal forest (Kulmala et al., 2004). More recently, Hallar et al. (2011) showed frequent NPF, correlated with UV irradiance, in background free tropospheric air. Studies have shown that the number concentrations of freshly nucleated particles at the start of NPF typically are in the range of $10^{3}$–$10^{4}$ cm$^{-3}$, but under certain conditions, for example in SO$_2$-enriched plumes and coastal areas, they can be higher than $10^{4}$ cm$^{-3}$ (Kulmala et al., 2004; Holmes, 2007; Bzdek and Johnston, 2010). Its spatial scale in regional events can extend beyond tens and up to several hundred kilometers (Stanier et al., 2004; Salma et al., 2011). The large scale of NPF have been recognized as a potentially significant contributor to the global CCN production (Merikanto et al., 2009). Most NPF observations have so far been made at mid and high-latitude regions of Northern Hemisphere (Kulmala et al., 2004) and the measurements in subtropical and tropical regions are sparse.

The favorable conditions and mechanisms leading to NPF and growth have been proposed for various atmospheric conditions. Kulmala et al. (2004), Bzdek and Johnston (2010), Hegg and Baker (2009), Holmes (2007), and Kulmala and Kerminen (2008) have illustrated the dynamic relationships among the nucleating (condensable) vapors, pre-existing aerosol densities, particle growth rates and nucleation rates. These studies also showed that NPF occurs when there are excess condensable vapors for the freshly nucleated particles to grow large without being scavenged by pre-existing aerosols. Zhang et al. (2012) reviewed recent advances in nucleation and growth and provided insights of these processes at the molecular level. Due to its low volatility at typical atmospheric temperatures, gas-phase sulfuric acid (H$_2$SO$_4$) is the most important nucleating vapor (Kuang et al., 2008; Kerminen et al., 2010; Sipila et al., 2010). Ternary species such as ammonia (NH$_3$), amines and other organic compounds can also enhance the nucleation rate by stabilizing the critical clusters (Benson et al., 2011; Erupe et al., 2011; Yu et al., 2012; Zhang et al., 2004).

Growth of new particles is governed by coagulation, condensation of low volatility and semi-volatile vapors, and heterogeneous reactions. Coagulation of freshly nucleated particles decreases the particle number and increases the modal size. However, its effect on particle growth is less important unless the number concentration and nucleation rate exceeds ~$10^5$ cm$^{-3}$ and $10^4$ cm$^{-3}$ s$^{-1}$, respectively (Anttila et al., 2010; Lehtinen et al., 2007). And, such conditions are rather uncommon but have been observed in polluted urban air, engine exhaust, industrial plumes or coastal areas. Yue et al. (2010) showed that intra- and extra-modal coagulation contributed 34 ± 17 % to particle growth during events with formation rates of 2–13 cm$^{-3}$ s$^{-1}$ in Beijing. Because condensation growth is strongly suppressed owing to the Kelvin effect for small particles, only vapors of low vapor pressure contribute to growth of freshly nucleated particles. Condensation of H$_2$SO$_4$ and subsequent stabilization by H$_2$O is considered irreversible under typical atmospheric conditions, whereas NH$_3$ plays a negligible role in nanoparticle growth (Zhang et al., 2009). However, field studies have found that H$_2$SO$_4$ condensation alone often fails to fully account for the observed particle growth (Stolzenburg et al., 2005; Yue et al., 2010), and it is possible that other semi-volatile organic vapors contribute to growth of new particles (Laaksonen et al., 2008; Zhang et al., 2009; Riipinen et al., 2011; Pierce et al., 2012). The Kelvin-effect barrier can be overcome by heterogeneous reactions, which can facilitate subsequent condensation of organic vapors (Zhang and Wexler, 2002). Also, amines and particulate-phase H$_2$SO$_4$ or organic acids lead to the formation of low volatility aminium salts that contribute to particle growth (Smith et al., 2010; Zhang et al., 2012).

Particle shrinkage can also occur due to evaporation of condensed species if the condensation or chemical reactions involved in particle growth are reversible. Suggested reversible processes with regard to new particle growth include the formation of aldol products and glyoxal oligomers in H$_2$SO$_4$ particles, and the physical adsorption of alcohols by H$_2$SO$_4$ solution (Zhang et al., 2012). In addition, particle-phase oxidation may result in enhanced volatilization of organic aerosols (Kroll et al., 2006; Molina, 2004). Modeling studies showed that vehicle-emitted particles smaller than 40 nm can shrink when the gas-phase concentration decreases upon dilution (Zhang and Wexler, 2004; Jacobson et al., 2005). In Hong Kong urban air, Yao et al. (2010) observed particle shrinkage following several NPF events and suggested that the particle shrinkage was related to the evaporation of organic components and ammonium nitrate. Recently, Backman et al. (2012) reported similar observations.
in São Paulo, Brazil, and suggested that the decrease of vapor concentrations by changes in atmospheric conditions could occur more rapidly than the decrease of vapor pressure by photochemical oxidation. Therefore, the net effect is the evaporation of the condensed phase to the gas phase.

Previously, Young et al. (2012) have characterized the spatiotemporal variability of 10–1000 nm submicrometer particles in various subtropical environments (urban, coastal, mountain and downwind) in Taiwan. They found that the variability of particle number is largely driven by primary emissions of local traffic and, to a lesser extent, by secondary NPF. Unlike the primary pathway, the secondary NPF is oftentimes accompanied with intense new particle growth over extended period of time. This NPF and growth are crucial processes for atmospheric aerosols because they can cause the new particles to grow to larger sizes of environmental health implications. With that in mind and building upon our previous work, here we aim to characterize the aerosol kinetics and properties during NPF and growth events and, along with meteorological conditions, to identify favorable conditions leading to those events. Also presented are several notable observations of new particle shrinkage following their prior formation and growth.

2 Experimental

2.1 Observation sites

The observation sites, instruments, meteorological conditions and data reduction methods are described in detail by Young et al. (2012). Observations were made at four air quality monitoring sites in the central Taiwan, a subtropical island country situated on the western edge of the Pacific Ocean and off the southeastern coast of mainland China. The sites are located 20–50 km apart in the Central Taiwan Air Quality Management District (CTAQMD) of ~7400 km² and nominally represent urban, coastal, mountain and downwind area (Fig. 1).

The urban site is in the center of the Taichung City basin, which has approximately 2610000 registered vehicles, ranking the 3rd in the country (Taiwan Ministry of Transportation and Communications, http://www.motc.gov.tw/). Among these vehicles, 65% are 4-stroke and 2-stroke scooters (ratio = 7 : 3), 30% are mostly gasoline-powered passenger cars, and the remaining 5% are trucks, buses and other types of vehicles. The coastal site is situated 4.8 km to the western coastline of Taiwan and 9 km south to the largest coal-fired power plant in Taiwan, which is also the world largest CO₂ emitting power plant (http://carma.org). Its annual emission of SO₂ and NOx in 2010 was 11,564 and 19,143 tons, respectively (http://www.epb.taichung.gov.tw/). The mountain site is located in a relatively rural basin; its east is the Central Mountain Ranges of height 2–3 km. The downwind site is near the southern border of the CTAQMD, 44 km to the southeast of the Taichung urban area. The prevailing northeasterly winds during winter months were conducive for transport of air pollutants from the upwind urban area to the downwind site.

The sampling campaigns were conducted consecutively at these four sites from October 2008 to January 2009 (cold season) and from August 2010 to October 2010 (warm season). The major meteorological differences between the cold and warm season include the average ambient temperature (20.6 vs. 27.6°C), prevailing wind direction (NNW vs. W) and accumulated precipitation (57.4 vs. 283.6 mm). The air quality was considered poorer during the cold season than the warm season, especially the all-sites hourly-average particle matter larger than 10 micrometer (PM₁₀, 57.1 µg m⁻³ during the warm season vs. 37.0 µg m⁻³ during the cold season), particulate matter larger than 2.5 micrometer (PM₂.5, 37.2 vs. 24.2 µg m⁻³), NOx (23.9 vs. 15.4 ppb) and SO₂ (3.8 vs. 2.9 ppb). During the entire study period, we collected a total 26075 aerosol number-size distributions over 137 effective sampling days (with the data coverage of ~92.5%).

2.2 Sampling and instrumentation

The aerosol measurement systems at all four observation sites were nearly identical and housed inside air-conditioned trailers. Ambient air was drawn through the site rooftop with a flow rate of 8 L min⁻¹ (lpm) via a 3-m vertical glass tube (ID = 2 cm, outside diameter OD = 2.54 cm). The outdoor section of the glass tube (2.5 m long) was shielded inside an aluminum tube. A short stainless steel forward-facing sampling probe was inserted into the indoor section of the
remaining glass tube (0.5 m). A small portion of the sample air (0.3 lpm) was then taken into the scanning mobility particle sizer (SMPS) through the sampling probe with a conductive tubing of 1.5-m maximum length. The diffusion loss to the glass and conductive tubing was estimated to be 15 % for the smallest measureable particles of 11.1 nm and 6 % for 25 nm particles, based on Baron and Willeke (2001). However, no loss corrections were made to the measured data because these transport losses were relatively small compared to the temporal variability of 59–195 % for the 10–25 nm particle number concentrations (Young et al., 2012). The air conditioners inside the trailers were activated only when the indoor temperature exceeded 31 °C. As a result, the sample air temperature basically followed the ambient air temperature. There were no water condensation issues with the sampling tube.

The particle number size distributions were measured with SMPS (GRIMM Aerosol Technik, Gmbh, Germany, Model 5.500) (e.g., Heim et al., 2004). The SMPS consists of an Am-241 neutralizer (Model 5.522), a long Vienna-type differential mobility analyzer (L-DMA; Model 55-900) and a butanol-based condensation particle counter (CPC; Model 5.403). With the default set up, the inner electrode of the L-DMA is positively charged. The CPC has a 50 % counting efficiency for particles as small as 4.5 nm, with the saturator and condenser temperature set at 40 °C and 15 °C, respectively. The SMPS was set to down-scan 6 min and 26 s from 10 000 V to 5 V plus a wait-time of 34 s, thus producing one average particle number size distribution every 7 min. The detectable aerosol mobility diameters range from 11.1 to 1083.3 nm (44 size bins) with a sheath and sample flow of 3 lpm and 0.3 lpm, respectively. The sheath airflow (3 lpm) was dried with a silica gel canister to a RH of ~10 % and then filtered through a HEPA before entering the differential mobility analyzer (DMA). The aerosol size distributions were therefore measured under dry conditions, and hygroscopic growth was not considered. For simplicity, we hereafter refer to the measureable size range as 10–1000 nm (as opposed to 11.1 to 1083.3 nm).

The size distributions inversion was carried out using the manufacturer-provided software (GRIMM 5.477 Version 1.35 Build 1), developed and described in detail by Riechsl (1991). In the software, the DMA and CPC correction efficiencies were taken into account in the calculations, whereas the charging correction is only available for single channel measurements and therefore not applied to the size distribution (i.e., multiple channels) measurements. To minimize multiple charging from large particles, an impactor with a nozzle diameter of 0.6 mm and 50 % size-cut of 1082 nm was attached to the inlet of the neutralizer and DMA. The SMPS has been routinely sent back to the manufacturer for calibration each year. In addition, the sizing accuracy was determined before the study by classifying NaCl of known mobility sizes (50, 76, 113, 168, and 241 nm) from a monodisperse aerosol generator (TSI; Model 3475), simultaneously with another SMPS (TSI; Model 3936). The differences between the expected and the measured particle sizes were less than 3 nm. The aerosol size distribution data were then subjected to quality control and assurance procedures proposed by Yu et al. (2004) that screen for abnormal size distributions. Only less than 0.1 % of aerosol data were identified as outliers and not included in the analysis. Following, the size-segregated particle number concentrations were derived by integrating the number concentrations of particles between 10–25, 25–100, 100–1000 nm. These concentrations hereafter are denoted as \( N_{10–25} \), \( N_{25–100} \), \( N_{100–1000} \), respectively. Hourly aerosol data were then computed to synchronize with the gas pollutants and meteorological data.

On-site air quality and meteorological data were obtained from the monitoring sites and validated by the Taiwan Environmental Protection Administration (TW EPA, 2009). At the sites, PM\(_{10}\) and PM\(_{2.5}\) were continuously monitored by means of \( \beta \)-ray attenuation (Met One Instruments, Oregon, USA, Model BAM-1020), SO\(_2\) by UV fluorescence (Ecotech, Victoria, Australia, Model 9850), NO\(_x\) by chemiluminescence (Ecotech, Victoria, Australia, Model 9841), CO by nondispersive IR (Horiba, Kyoto, Japan), and O\(_3\) by UV absorption (Ecotech, Victoria, Australia, Model 9810). Meteorological parameters used in the present study include temperature, relative humidity (RH), wind speed, wind direction and UV index (UVI). The UVI (1 UVI = 25 mW m\(^{-2}\)) was used in the calculation of UVB intensity. The evolution of the mixing height was estimated using the web-based Real-time Environmental Applications and Display sYstem (READY) from the Air Resources Laboratory of NOAA (http://ready.arl.noaa.gov/READYmet.php). The archive data were from the Global Data Assimilation System of National Centers for Environmental Prediction, which has a temporal and spatial resolution of 3 h and 1° of latitude and longitude, respectively.

### 2.3 NPF, growth and shrinkage event identification

A comprehensive review of the present instrumentation, best practices and other tools for studying atmospheric nucleation and NPF has recently been presented by Kulmala et al. (2012). In this study, the identification of NPF events, as well as growth and shrinkage, was made visually on a daily basis according to a set of distinct features in the temporal evolution of the aerosol number size distributions (e.g., Dal Maso et al., 2005; Stanier et al., 2004). An NPF event is characterized by a substantial increase of \( N_{10–25} \) followed by intense particle growth over a course over 1.5 h, in which the nucleation and growth rates can be determined with a good confidence level. The start and end time of NPF \( t_{\text{start}} \) and \( t_{\text{end}} \) is defined as the time when the \( N_{10–25} \) begin to rise sharply and when the particle growth stops, respectively. The \( t_{\text{start}} \) is an important parameter as it indirectly relates to the specific atmospheric conditions (such as mixing
height, temperature, and RH) and the traffic emission source strength, both of which exhibit strong temporal dependence. Particle shrinkage refers to the continuous decrease of particle sizes over \( \sim 2 \) h following their prior formation and growth.

### 2.4 Determination of particle formation, nucleation and growth rate

The empirical formation and nucleation rates \((J)\) during the NPF events were determined by a simplified approximation and an inverse modeling, respectively. The two methods are based on the general dynamic equation (GDE) in describing the evolution of aerosol size distribution (Seinfeld and Pandis, 2006). According to Dal Maso et al. (2005), the simplified GDE method determines the formation rate of 10 nm particles as follows:

\[
J_d = \frac{dN_{10-25}}{dt} + F_{\text{coag}} + F_{\text{growth}}
\]

where \( N_{10-25} \) is the number concentration of nucleated particles, \( F_{\text{coag}} \) is the loss of particles due to coagulation, and \( F_{\text{growth}} \) is the flux of particles out of the nucleated particle size range. The smallest measurable particle size was \( \sim 10 \) nm in this study, while the upper size range of the nucleated particle was set at 25 nm. Hence, the formation rate of 10 nm particles is hereafter denoted as \( J_{10} \). Assuming the homogeneous air mass with a constant source rate, \( dN_{10-25}/dt \) was determined by a linear fit between \( N_{10-25} \) and \( t \). In the same fashion, the modal growth rate (GR_{MOD}) was determined between the particle mode diameter (\( D_{\text{mode}} \)) and \( t \). \( F_{\text{coag}} \) was computed from (Dal Maso et al., 2005):

\[
F_{\text{coag}} = N_{10-25} \sum_j K_{ij} N_j
\]

where the summation of \( K_{ij} N_j \) is also known as the coagulation sink. \( N_j \) is the number concentration of size bin \( j \). \( K_{ij} \) is the coagulation coefficient between size bin \( i \) (i.e., the reference) and \( j \), and takes the below form (Seinfeld and Pandis, 2006):

\[
K_{ij} = 2\pi (d_i + d_j) (D_i + D_j) \beta_F
\]

where \( d \) is the midpoint diameter of a size bin, \( D \) and \( \beta_F \) are the size-dependent diffusion coefficient and Fuchs correction factor, respectively. It is noted that particle loss due to self-coagulation was neglected because of its minor contribution to NPF and growth, compared with inter-coagulation and condensation, especially in polluted environments (Anttila et al., 2010). \( F_{\text{growth}} \) was also neglected for simplicity, but in the present study the particles appeared to grow relatively fast (e.g., to above 25 nm within the first several hours). As a result, the above method underestimated the actual \( J_{10} \) because the particle transport losses and \( F_{\text{growth}} \) were both neglected.

The inversion model known as Particle Growth and Nucleation (PARGAN) has been used here to estimate the \( J \) and to supplement the \( J_{10} \), and its development and application have been described elsewhere (Verheggen and Mozurkewich, 2006; Erpe et al., 2010; Metzger et al., 2010). The online version of PARGAN is available at http://www.personal.kent.edu/~slee19/softwares.htm. In brief, the model implements a non-linear regression analysis of the GDE to fit the temporal evolution of the measured particle number size distributions. As such, the growth rate (GR_{PARGAN}) as a function of time can be determined and then used to estimate the time when the nucleation begins. Such highly time-resolved approach is therefore sensitive to any noise or variability in the measured size distributions. In this work, we assume the critical cluster has a diameter of 1 nm (Sipila et al., 2010). The \( J_1 \) is then derived by integrating particle losses from the time of nucleation to the time of measurement. This approach thus differs from others, because the GR_{PARGAN} is determined using a range of size intervals instead of total or modal particle number concentration and the \( J_1 \) is calculated solely on the basis of the measured size distributions, hence independent of any nucleation theories or parameterizations. The uncertainty therefore depends largely on the amount of extrapolation necessary to derive the correction factors for size-dependent particle number reductions (e.g., coagulation and dilution). These correction factors in turn depend on the assumed critical cluster size and the measured size distributions. Thus, erroneous assumption about the critical size and “noisy” measured size distributions will lead to increased uncertainty. In an atmospheric NPF study (Verheggen and Mozurkewich, 2002) and a smog chamber study (Verheggen and Mozurkewich, 2006), the authors estimated that the overall uncertainty of the PARGAN derived \( J \) to be a factor of 2 to 3, which is within an acceptable order-of-magnitude limit reported for other NPF analysis tools by Korhonen et al. (2011). It is nevertheless recommended that the interpretation and comparisons of modeling results should be proceeded with cautions. Unless equipped with instruments capable of detecting \( \sim 1 \) nm particles, the highly time-resolved PARGAN approach still provides useful information and reference about the mechanism and intensity of atmospheric nucleation.

### 2.5 Calculation of condensation sink, condensable vapour and H2SO4 proxy concentration

Condensation sink (CS) is a measure of loss rate of H2SO4 vapor on pre-existing particles and is proportional to the surface area density of aerosol particles. Elevated pre-existing particles can suppress aerosol nucleation by scavenging freshly nucleated particles and condensable vapors. Assuming H2SO4 as the condensable species, the CS in the transition regime was calculated as follows (Kulmala et al., 2001):

\[
CS = 2\pi D_0 \sum_j d_j \beta_{m,j} N_j
\]
where $D$ is the diffusion coefficient of H$_2$SO$_4$ (0.104 cm$^2$ s$^{-1}$) estimated by the FSG-LaBas method (Lyman et al., 1990), and $\beta_m$ is the size dependent transitional correction factor (Fuchs and Sutugin, 1971).

Based on the observed growth rates and assuming the condensable vapor was solely H$_2$SO$_4$, the condensable vapor concentration ($C$) can be estimated according to Kulmala et al. (2005) as follows:

$$C = \frac{\rho}{\Delta t Dm_v} \left( \frac{d^2_p - d^2_{p,0}}{8} + \left( 2 \frac{2}{3\alpha} - 0.312 \right) \lambda (d_p - d_{p,0}) \right) + 0.623 \lambda^2 \ln \left( \frac{2\lambda + d_p}{2\lambda + d_{p,0}} \right)$$

(5)

where $\rho$ is the particle density (1.84 g cm$^{-3}$), $D$ is the diffusion coefficient (0.104 cm$^2$ s$^{-1}$), $m_v$ is the molecular mass of the condensable vapor (1.63 $\times$ 10$^{-22}$ g molec$^{-1}$), $\alpha$ is the mass accommodation coefficient (1), and $\lambda$ is the mean free path (6.7 $\times$ 10$^{-6}$ cm) at 1 atm and 298 K.

Ambient gas phase H$_2$SO$_4$ is known to be involved in the NPF and their subsequent growth. Its production and concentration depend on the precursor SO$_2$ and OH concentrations, between which the latter is dependent on photolysis rate of O$_3$ by solar radiation in the 280–320 nm UVB range. Its major sink is condensation to pre-existing aerosols. In the cases where direct measurements are unavailable, a few methods have been proposed to estimate H$_2$SO$_4$ proxy concentrations. Here, we have adopted the proxy estimation method of H$_2$SO$_4$ concentration ($P$) by Petäjä et al. (2009) as follows:

$$P = k_0' = k \frac{[SO_2]UVB}{CS}$$

(6)

where $k'$ is the scaling factor and CS is the condensation sink. The SO$_2$ concentration has unit of cm$^{-3}$. In this study, the UVB was determined from the measured daily maximum UVI, where 1 UVI equals to 25 mW m$^{-2}$ (or 90 J m$^{-2}$).

### 3 Results and discussion

#### 3.1 Characteristic time of NPF events in Taiwan

Among the 137 sampling days at the four observation sites, we identified 14 intense NPF events. The NPF and growth date, time, duration and daily meteorological conditions are summarized in Table 1. As shown, 11 of them took place in the warm season (August–September) and three in the cold season (December–January). In addition, nine events occurred at the urban site, four at the downwind site, one at the coastal site, and none at the mountain site. The mountain site had the highest average PM$_{2.5}$ of 47.7 and 36.9 µg m$^{-3}$, whereas the PM$_{10}$ were 66.7 and 49.3 µg m$^{-3}$ in the cold and warm study season, respectively (Young et al., 2012). The high PM levels thus may have suppressed the NPF. All the events, except one (17 December 2008), started before ~10:30 LT and ended before ~15:00 LT. These $t_{\text{start}}$ and $t_{\text{end}}$ are similar to those observed in Beijing, China (Wu et al., 2007), Pittsburgh, PA, USA (Stanier et al., 2004) and Hyytiälä, Finland (Boy and Kulmala, 2002). Nevertheless, some NPF events were limited to morning hours, whereas others occurred during midday hours. The duration ($\Delta t$) from the start of NPF to the end of particle growth lasted for 1.5 to 5 h. It is notable that eight NPF events took place over consecutive days (10–17 August 2010) at the urban site in the warm season (Fig. A1). Furthermore, in five of the 14 NPF events, the growth was immediately followed by intense particle shrinkage. Four shrinkage events were observed in the warm season with two at the urban site and the other two at the downwind site, and only one in the cold season at the coastal site. On event days, the ambient temperature and RH were below 18°C and 67% in the cold season, whereas mostly above 27°C and 69% in the warm season. The wind speeds were considerable stagnant (mostly $<1.3$ m s$^{-1}$) during all the event days except the one at the coastal site. The prevailing winds were from the northeast in the cold season whereas from the south or east in the warm season.

#### 3.2 Air pollutants, condensable vapor and H$_2$SO$_4$ proxy concentrations

The daily PM$_{2.5}$, CS, SO$_2$, condensable vapor and H$_2$SO$_4$ proxy concentrations during event days are given in Table 2. As shown, the PM$_{2.5}$ were typically below 20 µg m$^{-3}$, substantially lower than the annual average value of the hourly PM$_{2.5}$ (35.8 µg m$^{-3}$). Only three event days had PM$_{2.5}$ above 26.4 µg m$^{-3}$ and up to 46.6 µg m$^{-3}$. The daily CS and SO$_2$ were in the range of 1.6–3.9 $\times$ 10$^{-2}$ s$^{-1}$ and 2.0–4.6 ppb, respectively. The above PM$_{2.5}$ and CS on event days were similar to, whereas the SO$_2$ were lower than those observed in Pittsburg, PA (Stanier et al., 2004) and Beijing, China (Wu et al., 2007). Figure 2 compares the daily CS and SO$_2$ between non-event and event days. As shown, both the CS and SO$_2$ on event days were typically below the median values on non-event days; for reference, the median CS for the entire study period is 2.6 $\times$ 10$^{-2}$ s$^{-1}$. This indicates that the NPF events were favored on days with low concentrations of both pre-existing particles and SO$_2$. The preference for low SO$_2$ is somewhat counterintuitive as it is the precursor of the major nucleating vapor H$_2$SO$_4$. Nevertheless, the lower or comparable SO$_2$ on event days than on non-event days suggests that SO$_2$ was not a limiting factor for NPF. For example, Salma et al. (2011) found that the SO$_2$ (yearly median 2.6 ppb) is always in excess for NPF in Budapest, Hungary. In this study, there are several possible reasons to the rather counterintuitive preference for low SO$_2$: first, the minimum SO$_2$ of ~2.0 ppb at the observation sites is already sufficient for NPF. Second, the rate-limiting step in the production of H$_2$SO$_4$ involves the oxidation of SO$_2$ by OH. Therefore, elevated SO$_2$ alone does not necessarily warrant
Table 1. The NPF and growth event date, time, duration, and daily averages of ambient temperature ($T$), relative humidity (RH), wind speed, and prevailing wind direction at the observation sites.

<table>
<thead>
<tr>
<th>Date</th>
<th>Site</th>
<th>Day of the Week</th>
<th>Shrink</th>
<th>Start/End time</th>
<th>$T$ ($^\circ$C)</th>
<th>RH (%)</th>
<th>Wind Speed (m s$^{-1}$)</th>
<th>Prevailing Wind Direction</th>
</tr>
</thead>
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</table>

1 The four observation sites: U indicates urban, C coastal and D downwind site. There were no NPF events observed at the mountain site.  
2 “Yes” for days with particle shrinkage following their prior formation and growth.  
3 The duration between the start of NPF to the end of particle growth.
periods given by Young et al. (2012). The average \( N_{10–25} \) during the NPF and growth events were in the range of 1.2–4.0 \( \times 10^4 \) cm\(^{-3} \), which were \( \sim 1.7 \) to 5.7 times higher than average levels of 0.7 \( \times 10^4 \) cm\(^{-3} \). The average \( N_{25–100} \) were in the range of 1.3–3.3 \( \times 10^4 \) cm\(^{-3} \), which were \( \sim 1.2 \) to 3.0 times higher than average levels of 1.1 \( \times 10^4 \) cm\(^{-3} \). It is clear that the NPF events resulted in substantial increases of the \( N_{10–100} \) (i.e., UFPs). As shown in the hourly data (Fig. A2), the peak \( N_{10–100} \) at 10:00 LT on event days were preceded by a decrease of CS, coincided with a slight increase of \( \text{SO}_4 \) and strong atmospheric dilution, and lagged 2 h behind the traffic CO and \( \text{NO}_x \) peaks. These results suggest that the elevated \( N_{10–100} \) at 10:00 LT were not emitted directly from motor vehicles but formed shortly afterwards in the atmosphere. On non-event days, the diurnal variations of \( N_{10–100} \) were nearly identical to those of \( \text{NO}_x \) and CO (Fig. A2). This indicates that the particle number variability on non-event days was largely driven by traffic emissions. Similar traffic-related \( N_{10–100} \) peaks (i.e., ones that occurred at \( \sim 08:00 \) LT) were also observed at a downwind site in urban Budapest, Hungary (Salma et al., 2011). Although less common, the average \( N_{10–100} \) of 7.5 \( \times 10^4 \) cm\(^{-3} \) during NPF was \( \sim 3.6 \) times higher than that of 2.1 \( \times 10^4 \) cm\(^{-3} \) emitted from the morning traffic. The approximation-derived formation rates of 10 nm particles \( (J_{10}) \) during NPF events at the urban site were 7.0–30 cm\(^{-3} \) s\(^{-1} \), whereas at the coastal and downwind sites they were 4.4–15 cm\(^{-3} \) s\(^{-1} \). The coagulation loss \( (F_{\text{coag}} \) in Eq. 1) of the newly-formed 10–25 nm particles at the urban site accounted for 45–80 % of the derived \( J_{10} \), whereas at the coastal and downwind sites its contributions to \( J_{10} \) were more variable, between 28–85 %. Overall, these losses on average accounted for 60 ± 17 % of the \( J_{10} \). This shows that the apparent particle flux \( (dN/dt) \) alone would have largely underestimated the \( J_{10} \). The observed \( J_{10} \) in this study are generally higher than that reported in the studies over a large number of measurements sites worldwide summarized by Kulmala et al. (2004). Recent studies showed that the \( J_1 \) in Hong Kong, Beijing, Shangdianzi and New Delhi were 1.8–5.1 cm\(^{-3} \) s\(^{-1} \), 2.0–81 cm\(^{-3} \) s\(^{-1} \), 0.7–73 cm\(^{-3} \) s\(^{-1} \) and 3.3–14 cm\(^{-3} \) s\(^{-1} \), respectively (Mönkkönen et al., 2005; Shen et al., 2011; Wu et al., 2007; Yao et al., 2010; Yue et al., 2010).

The PARGAN-derived nucleation rates \( (J_1) \) during NPF events at the urban site were in the range of 132–253 cm\(^{-3} \) s\(^{-1} \), whereas at the coastal and downwind sites they were in the range of 40–113 cm\(^{-3} \) s\(^{-1} \) (Table 2). The observed \( J_1 \) are in line with the modeled median nucleation rates of 311 cm\(^{-3} \) s\(^{-1} \) and 18 cm\(^{-3} \) s\(^{-1} \) in polluted boundary layer and the rural/continental background, respectively (Aarnio et al., 2010; Kulmala et al., 2004), yet substantially lower than the nucleation rates of 3 \( \times 10^5–5 \) \( \times 10^7 \) cm\(^{-3} \) s\(^{-1} \) in coastal areas (Pirjola et al., 2002). Using the same PARGAN method (Erupse et al., 2010; Kanawade et al., 2012), the median \( J_1 \) of \( \sim 11 \) cm\(^{-3} \) s\(^{-1} \) in semi-rural Kent, OH, USA, was substantially lower than the present study. This difference may be due to the substantially higher \( \text{SO}_2 \) and other possible precursor nucleating vapors at the present observation sites. On the other hand, based on the growth rates and Fuchs surface area, Iida et al. (2008) estimated the \( J_1 \) in Mexico City and New Delhi were 1900–3000 cm\(^{-3} \) s\(^{-1} \) and \( \sim 1300 \) cm\(^{-3} \) s\(^{-1} \), respectively. These values are substantially higher than the \( J_1 \) in this study, possibly because

<table>
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<th>Date</th>
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<th>Shrink2</th>
<th>PM(_{2.5}) ((\mu)g m(^{-3}))</th>
<th>CS (10(^{-2}) cm(^{-3}) s(^{-1}))</th>
<th>SO(_2) (ppb)</th>
<th>C (10(^8) cm(^{-3}))</th>
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1. The four observation sites: U indicates urban, C coastal and D downwind site. There were no NPF events observed at the mountain site.
2. “Yes” for days with particle shrinkage following their prior formation and growth.
3. The \( \text{H}_2\text{SO}_4 \) proxy concentration assuming the scaling factor \( s \) of 9.9 \( \times 10^{-4} \) m\(^2\) W\(^{-1}\) s\(^{-1}\) reported by Petaja et al. (2009).

Table 2. The daily averages of PM\(_{2.5}\), condensation sink (CS), \( \text{SO}_2 \), condensable vapor (C) and \( \text{H}_2\text{SO}_4 \) proxy (P) concentrations during the NPF event days.
the two megacities are substantially more polluted with respect to SO$_2$ and PM than the present study area (Gurjar et al., 2008; Raga et al., 2001).

The PARGAN-derived growth rates ($\text{GR}_{\text{PARGAN}}$) and approximation-derived GR$_{\text{MOD}}$ at the urban site were 6.5–9.5 nm h$^{-1}$ and 6.7–24 nm h$^{-1}$, respectively. At the coastal and downwind sites, they were 6.7–15 nm h$^{-1}$ and 7.4–15 nm h$^{-1}$, respectively. The percentages of GR$_{\text{PARGAN}}$ over GR$_{\text{MOD}}$ ranged from 51% to 187%, with an average of 88±41%, similar to the percentage shown by Erupe et al. (2010). Thus, the differences between the GR$_{\text{PARGAN}}$ and GR$_{\text{MOD}}$ are within a factor of 2, indicating reasonable consistency between the two different approaches to GR estimations.

The GR$_{\text{MOD}}$ during particle shrinkage in this study were between −5.1 and −7.6 nm h$^{-1}$, which were slightly smaller than that between −8.2 and −10.7 nm h$^{-1}$ observed in Hong Kong (Yao et al., 2010). The ratio of shrinkage-to-growth rates for each shrinkage event day was in the range of 0.40–0.66, with an exception for the shrinkage event on the 7 September 2010 during which the ratio was 1.02. The evaporation loss of prior condensed vapor can be estimated by the particle volume change as a result of diameter shrinkage. The results show that on the 3 January 2009 and the 7 September 2010 the volume losses were above 90%, suggesting nearly complete evaporation. On other remaining days, the volume losses ranged from 18% to 38%. These results indicate that a notable fraction of the originally condensed chemical species evaporated from the particle phase to the gas phase under atmospheric conditions. Bzdek et al. (2012) recently showed that sulfate contributed to 41–46% new particle growth, whereas NO$_2^+$, NH$_4^+$ and organics contributed to the remaining 54–59% growth. In particular, some of the latter chemical species are known to be semi-volatile and, as such, they have the potentials to evaporate off the new particles under favorable conditions.

A closer examination of the size distribution data reveals that the two shrinkage events (3 January 2009 and 5 September 2010) showed simultaneous decrease of particle size and $N_{10−25}$, similar to the Hong Kong shrinkage events (Yao et al., 2010). However, the other three shrinkage events (12 August, 16 August and 7 September 2010) showed decreasing particle size with increasing $N_{10−25}$. Furthermore, the grown particles in some cases can shrink back to the smallest measurable size of ~10 nm, whereas a few others are similar to the Hong Kong shrinkage events where the particle shrinking stopped at sizes above the smallest measurable size at ~20 nm in our study (Fig. A3). As a result, the new particle growth and shrinkage creates a unique “arch-like” shape in the size distribution contour plot.

### 3.4 Case studies of NPF and growth

As described earlier in Sect. 3.1, all but one NPF events started before 10:30 LT. Nevertheless, some events were more or less limited to morning hours, whereas others lasted into midday hours. Two typical examples of NPF and growth events are given below, one during morning hours and the other during midday hours. On the 17 August 2010, the $N_{10−25}$, CS, CO and NO$_x$ started to increase from 05:00 LT and then reached a maximum at ~08:00 LT due to the increased traffic emissions during rush hours (Fig. 3).

---

### Table 3. The size-segregated particle number concentration, formation/nucleation and growth/shrinkage rates during the NPF events.

<table>
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<tr>
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<th>Formation/Nucleation Rate (cm$^{-3}$ s$^{-1}$)</th>
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1. The four observation sites: U indicates urban, C coastal and D downwind site. There were no NPF events observed at the mountain site.
2. “Yes” for days with particle shrinkage following their prior formation and growth.
3. The average concentrations from the start of NPF to the end of particle growth.
The traffic-emitted \( N_{10-25} \) at 08:00 LT had a maximum of \( 1.7 \times 10^4 \) cm\(^{-3} \). Between 08:00 and 09:00 LT, the size distribution contour plot shows a significant decrease of particles larger than 50 nm, indicating strong atmospheric dilution at that time. As a result, the primary pollutants CO, NO\(_x\), and CS decreased to a minimum at \( \sim 10:00 \) LT. At about the same time, the \( N_{10-25} \) sharply increased from \( 1.5 \times 10^4 \) cm\(^{-3} \) to \( 5.8 \times 10^4 \) cm\(^{-3} \) at 10:00 LT. The \( N_{10-25} \) peak coincided with a SO\(_2\) peak of 4.3 ppb, indicating possible involvement of H\(_2\)SO\(_4\) (as opposed to direct traffic emissions) in NPF. It is important to note that the SO\(_2\) peak lagged 2 h after the CO and NO\(_x\) peaks, thus suggesting the SO\(_2\) was not directly emitted from motor vehicles. The high level of \( N_{10-25} \) was then followed by intense particle growth from 11 to 62 nm over a duration of \( \sim 4 \) h. The \( J_1 \) and GR\(_{PARGAN}\) during the first 2 h of the event varied around \( 200 \) cm\(^{-3}\)s\(^{-1}\) and 5 nm h\(^{-1}\), respectively. The GR\(_{PARGAN}\) then increased to \( \sim 15 \) nm h\(^{-1}\) in the final few hours. Prior to 10:00 LT, the relatively light winds came from the southeast under a cloudless and sunny condition (9.7 UVI). The O\(_3\) had a maximum of 65 ppb at 13:00 LT. Near the end of the NPF event, the winds became westerly and another \( N_{10-25} \) peak was observed at 15:00 LT. Unlike the morning NPF event, this afternoon particle formation event was related to plume because of the concurrently elevated primary pollutants SO\(_2\) (17 ppb), NO\(_x\) (36 ppb) and CO (0.5 ppm), and a decrease of O\(_3\) (37 ppb). In addition, there was no particle growth following the particle formation, possibly because of the extremely high CS (7.1 \( 10^{-2}\) s\(^{-1}\)) and decreased photochemical activities (low UVI and O\(_3\)).

As opposed to the previous morning NPF, the following example occurred during midday hours that are typically characterized by low concentrations of primary pollutants CO and NO\(_x\), but elevated O\(_3\) and mixing height (Fig. A2). On the 3 December 2008, the \( N_{10-25} \) and CO and NO\(_x\) started to increase from 05:00 LT and then reached a maximum at \( \sim 08:00 \) LT (Fig. 4). The traffic-emitted \( N_{10-25} \) at 08:00 LT had a maximum of \( 1.4 \times 10^4 \) cm\(^{-3} \), which is similar to that observed on other event days. After 08:00 LT, the CO and NO\(_x\) then decreased substantially due to decreased source strength and increased mixing height. Strong atmospheric dilution caused the CS to drop to a minimum of 2.2 \( 10^{-2}\) s\(^{-1}\) at 13:00 LT. At about the same time, the \( N_{10-25} \) increased sharply from \( 0.5 \times 10^4 \) cm\(^{-3} \) to \( 3.5 \times 10^4 \) cm\(^{-3} \) within an hour. The only pollutant that coincided with the \( N_{10-25} \) peak was the SO\(_2\) (5 ppb). Similar to the previous example, the SO\(_2\) peak lagged 4 h after the CO and NO\(_x\) peaks, suggesting the SO\(_2\) was not directly emitted from motor vehicles. The NPF event was then followed by intense particle growth from 11 to 39 nm during the following two hours. The \( J_1 \) reached as high as \( \sim 300 \) cm\(^{-3}\)s\(^{-1}\) at 12:30 LT, whereas the GR\(_{PARGAN}\) were in the range of 2–10 nm h\(^{-1}\). The south-easterly winds in the morning were stagnant (\(< 0.6 \) m s\(^{-1}\)). Afterwards, the wind direction changed to northerly and the wind speed increased significantly to above 1 m s\(^{-1}\). The UVI indicates it was a cloudless, sunny day. Despite the maximum UVI of 5.2 at 12:00 LT was relatively weak (compared to Fig. 3d), the O\(_3\) peak of 84 ppb at 16:00 LT was significantly higher than that at the urban site in the warm season (e.g., Fig. 3c; 65 ppb O\(_3\)). This suggests that photochemical activities could be quite strong at the downwind site in the cold season. The termination of the midday NPF event
Fig. 4. The same as Fig. 3, except this is for a midday NPF and growth event on the 3 December 2008.

at 16:00 LT coincided with elevated CS ($5.5 \times 10^{-2}$ s$^{-1}$) and SO$_2$ (6.2 ppb). This result suggests that the observation site was impacted by polluted plume near the end of NPF due to changing wind direction. There were elevated $N_{10-25}$ of $1.7 \times 10^4$ cm$^{-3}$ during the morning and evening traffic rush hours. Comparing the $N_{10-25}$ levels between the traffic at about ~08:00 LT and the NPF after ~09:00 LT, the intensity of UFP production via the secondary pathway was generally stronger than the primary pathway.

From above and other events in this study, we observed that the onset of morning or midday NPF events coincided with decreases of CS and increases of SO$_2$ during periods (09:00–13:00 LT) with enhanced atmospheric mixing and dilution. The decrease of CS, as well as traffic-emitted CO and NO$_x$, due to dilution is rather expected; however, the increase of SO$_2$ suggests it was not from local traffic emissions but likely from power plant or industrial plumes transported or mixed down from aloft to the observation sites. In the case of the morning events, the short time lag between the strong traffic emissions and the start of NPF before 09:30 LT indicates that the morning NPF may have been facilitated by the precursor nucleating/condensable vapors from traffic emissions.

3.5 Case studies of particle shrinkage events

Particle shrinkage events were identified during five out of the 14 NPF events. Two shrinkage events were observed at the urban site, two at the downwind site and one at the coastal site. In the following, we present two examples of particle shrinkage, one where the $D_{mode}$ and $N_{10-25}$ were positively correlated and the other negatively correlated. The evolution of particle number size distributions during the remaining three events can be found in Fig. A3.

On the 12 August 2010 at the urban site in the warm season, the start of morning NPF occurred at ~08:00 LT in the midst of traffic rush hours (Fig. 5). It was then immediately followed by intense particle growth from 11 to 51 nm until 12:30 LT. The maximum $N_{10-25}$ of $5.2 \times 10^4$ cm$^{-3}$ at 10:00 LT was preceded by a local minimum CS of $1.7 \times 10^{-2}$ s$^{-1}$, and it was correlated with a SO$_2$ peak of 3.6 ppb. The median $J_1$ and $GR_{PARTG}$ were 216 cm$^{-2}$ s$^{-1}$ and 7.4 nm h$^{-1}$, respectively. Between 12:30 and 16:00 LT, the $D_{mode}$ and CS decreased from 32 to 11 nm and $3.1 \times 10^{-2}$ to $1.5 \times 10^{-2}$ s$^{-1}$, respectively. It is notable that the shrinkage occurred during periods with the strongest atmospheric dilution (max. mixing height 590 m), highest ambient temperature (max. 34°C) and lowest relative humidity (min. 61%). The continuous decrease of CS as well as CO and NO$_x$ during the shrinkage period suggests that the site was not impacted by local primary emissions. Meanwhile, the $N_{10-25}$ and SO$_2$ were increasing with time. The $N_{10-25}$ reached another maximum of $5.1 \times 10^4$ cm$^{-3}$ near the end of the shrinkage. The $GR_{MOD}$ during the growth and shrinkage were 11 and $-5.6$ nm h$^{-1}$, respectively. The winds of below 1.7 m s$^{-1}$ were mild throughout the day. The winds were from the south before 12:00 LT, but then changed to westerly during the shrinkage period. The change of wind direction was accompanied with a slight increase of wind speed and the appearance of cloud clover, as indicated by the sudden drop of UVI at 14:00 LT. The maximum O$_3$ was 51 ppb. The SO$_2$ peak of 3.4 ppb during the shrinkage period was of particular interest as the rise of SO$_2$ levels was simultaneously observed at three other air quality monitoring sites.
near the urban basin (Fig. A4). These results suggest that the SO$_2$-enriched and low-CS air mass due to atmospheric advection and mixing may in turn have caused another NPF event near the end of particle shrinkage period, while the enhanced atmospheric dilution and hot/dry ambient conditions favor the evaporation of semi-volatile species from the particles to the gas phase. The former explains the increasing $N_{10-25}$ whereas the latter results in the decreasing $D_{mode}$.

On the 3 January 2009 at the coastal site in the cold season, the midday NPF between 10:30 and 13:00 LT was followed by intense particle shrinkage (Fig. 6). At the start of NPF, the $N_{10-25}$ increased rapidly from $1.0 \times 10^4$ cm$^{-3}$ to $2.5 \times 10^4$ cm$^{-3}$. The SO$_2$ also increased from 2.1 to 4.0 ppb. Afterwards, the particles grew from 11 to 36 nm. The median $J_1$ and GR$_{PARGAN}$ were 99 cm$^{-2}$ s$^{-1}$ and 15 nm h$^{-1}$, respectively. Then the $N_{10-25}$ gradually decreased to $0.6 \times 10^4$ cm$^{-3}$ by 19:00 LT and the $D_{mode}$ decreased from 30 to 11 nm between 14:00 and 17:00 LT. The GR$_{MOD}$ during the growth and shrinkage were 7.8 and $-5.1$ nm h$^{-1}$, respectively. The elevated CS, SO$_2$, CO and

Fig. 5. The same as Fig. 3, except this is for a NPF, growth and shrinkage event on the 12 August 2010.

Fig. 6. The same as Fig. 3, except this is for a NPF, growth and shrinkage event on the 3 January 2009.
NO$_3$ at 16:00 LT indicate an impact of polluted air. However, it did not affect the apparent particle shrinkage. Although not measured on-site, a nearby monitoring site (8.3 km to the southeast) shows that it was a cloudless, sunny day with the maximum UVII of 4.1 between 12:00 and 13:00 LT. The maximum O$_3$ was 47 ppb. The prevailing north-northeasterly winds were strong (> 4 m s$^{-1}$), cool and stable throughout the event day. It is also notable that the temperature increased from a minimum 12°C to a maximum 20°C between 12:00 and 16:00 LT. These results suggest that the strong dilution due to elevated wind speed and the large temperature rise were likely the driving forces leading to the evaporation of semi-volatile species from the particles to the gas phase. However, unlike the previous example, the above particle shrinkage is characterized by a simultaneous decrease of $D_{\text{mode}}$ and $N_{10-25}$.

The above observations clearly show that new particle growth can be a reversible process under certain atmospheric conditions. For example, enhanced dilution and hot/dry ambient conditions can decrease the ambient vapor partial pressure or increase the equilibrium vapor pressure over the particle surface. As such, the condensed vapors tend to evaporate off the particle phase. At present it is unclear what were the evaporating vapors involved during the particle shrinkage. Potential candidates include semi-volatile ammonium nitrate ($\text{NH}_4\text{NO}_3$) and organics compounds due to their high volatility and abundance in ambient air. In central Taiwan, Fang et al. (2006) showed that the major ionic species in ultrafine particles are SO$_4^{2-}$, followed by NH$_4^+$ and NO$_3^-$. Furthermore, Lin et al. (2006) reported the annual average of NH$_3$ was 12.3 ppb, with the highest average of 16.4 ppb observed in the summer. More recently, Bzdek et al. (2012) showed that 29–46 % of the total mass growth of new particles were attributable to SO$_4^{2-}$ at an urban site in Wilmington, DE, USA. The remaining, more than 50 %, new particle growth was due to NO$_3^-$, NH$_4^+$ and organics. In addition, Riipinen et al. (2012) suggested that the reversible net condensation of gas-phase oxidation products is an important process governing organic vapor uptake by nanoparticles. The above studies demonstrate the potentially abundant and semi-volatile NH$_4^+$, NO$_3^-$ and organics in the particle phase, and thus we speculate that they may have been involved in the particle shrinkage in the present study.

4 Conclusions

This study provides the first systematic analysis for new particle formation (NPF) and growth events at four distinct types of environment (urban, coastal, mountain and downwind) in an air quality management district of subtropical central Taiwan. A total of 14 NPF and growth events were identified from October 2008 to January 2009 (cold period) and from August 2010 to October 2010 (warm period). The events were then analyzed for aerosol dynamics and properties, air pollutant and meteorological conditions. Some NPF events were more or less limited to morning hours, whereas others occurred during midday hours. In either cases, the onset of NPF events coincided with decreases of CS and increases of SO$_2$ during periods (09:00–13:00 LT) under enhanced atmospheric mixing and dilution. The increase of SO$_2$ during enhanced dilution, in particular, suggests that the SO$_2$ was not from local traffic emissions but likely from power plant or industrial plumes transported or mixed down from aloft to the observation sites. Nevertheless, the lower or comparable SO$_2$ on event days than on non-event days suggests that SO$_2$ was not a limiting factor for NPF.

Overall, the NPF events contributed significantly to the ambient ultrafine particles, with the maximum concentrations of 2.5–8.9 × 10$^4$ cm$^{-3}$, which were 2–4 times higher than that due to typical traffic emissions. The approximation-derived particle formation rate $J_{10}$ and growth rates GR$_{\text{MOD}}$ were in the range of 4.4–30 cm$^{-3}$ s$^{-1}$ and 7.4–24 nm h$^{-1}$, respectively. The nucleation rates $J_1$ and growth rates GR$_{\text{PARGAN}}$ inverting from the measured aerosol size distributions were in the range of 40–253 cm$^{-3}$ s$^{-1}$ and 6.5–15 nm h$^{-1}$, respectively.

A unique feature of our observations is the particle shrinkage following their prior formation and growth. In the five shrinkage events, their respective GR$_{\text{MOD}}$ ranged from $-5.1$ to $-7.6$ nm h$^{-1}$. The shrinking and corresponding particle volume losses indicate that a notable fraction of the originally condensed chemical species was semi-volatile and had evaporated from the particle phase to the gas phase under atmospheric conditions. In intense cases where particles shrank back to the smallest measurable size of $\sim 10$ nm, the new particle growth and shrinkage thereby created a unique “arch-like” shape in the size distribution contour plot. Time-series analyses of air pollutants and meteorological conditions indicate that the particle shrinkage was related to air masses with enhanced atmospheric dilution, high ambient temperature and low relative humidity. Such atmospheric conditions favor the evaporation of semi-volatile species from the particle phase to the gas phase. Although chemical identification of the potential evaporating vapors was not possible in the present study, we speculate that the candidate vapors can be NH$_4^+$, NO$_3^-$ and organics due to their semi-volatile nature and abundance in the study area. Chemical and volatility analysis of the nucleating and condensing/evaporating vapors is required in the future to infer more definitive mechanisms of NPF, particle growth and shrinkage that are relevant to climate change and public health.

Appendix A

A1 Eight consecutive morning NPF events with two particle shrinkage events

Figure A1 shows the diurnal variations of the number size distributions during eight consecutive morning NPF events
Fig. A1. The measured particle number size distributions during eight consecutive NPF events at the urban site from the 10 to the 17 August 2010. Among them there were two particle shrinkage events on the 12 and the 16 August 2010.

Fig. A2. Diurnal variations of the average (a) mixing height, (b) NO\textsubscript{x}, (c) CO, (d) condensation sink (CS), (e) SO\textsubscript{2} and (f) 10–100 nm particle number (N\textsubscript{10–100}) concentrations on non-event and event days.
Fig. A1. The measured particle number size distributions, $D_{\text{mode}}$ (stars), $N_{10-25}$ (solid lines) and CS (dash lines) during the three other particle shrinkage events on the (a) 16 August, 2010, (b) 5 September 2010 and (c) 7 September 2010. As shown, the grown particles in the first event (a) shrank back to the smallest measurable particle size of $\sim 10$ nm, whereas in the remaining two events (b and c) they shrank back to sizes of $\sim 20$ nm and then stopped. The $N_{10-25}$ during the first and third events were increasing with time, whereas during the second event they were decreasing with time.

A2 Diurnal variations of mixing height, air pollutants, condensation sink and ultrafine particle number on non-event and event days

Figure A2 presents the diurnal variations of the average mixing height, air pollutants ($\text{NO}_x$, CO, $\text{SO}_2$), condensation sink (CS), and 10–100 nm ultrafine particle number ($N_{10-100}$) concentrations on non-event and event days. As shown, the differences of hourly pollutant levels and mixing height between the non-event and event days were not statistically significance. Nevertheless, the diurnal trends of the hourly-average CS and $N_{10-100}$ on events days deviated away from those on non-event days during late morning hours, between 09:00–12:00 LT. On event days, the 2 h lag between the $\text{NO}_x$ and CO peaks at 08:00 LT and the $N_{10-100}$ peak at 10:00 LT suggests that the elevated ultrafine particles was not emitted directly from motor vehicles but formed shortly afterwards in the atmosphere, i.e., via NPF. On non-events days, the diurnal variations of $N_{10-100}$ were nearly identical to those of $\text{NO}_x$ and CO, indicating strong influence from traffic emissions.

A3 Three other particle shrinkage events

Figure A3 shows the measured number size distributions, $D_{\text{mode}}$, $N_{10-25}$ and CS during the three other particle shrinkage events on the (a) 16 August 2010, (b) 5 September 2010 and (c) 7 September 2010. As shown, the grown particles in the first event (a) shrank back to the smallest measurable particle size of $\sim 10$ nm, whereas in the remaining two events (b and c) they shrank back to sizes of $\sim 20$ nm and then stopped. The $N_{10-25}$ during the first and third events were increasing with time, whereas during the second event they were decreasing with time.

A4 Regional-scale $\text{SO}_2$ impact

Figure A4 shows the diurnal variations of $\text{SO}_2$ at the present urban site (JM) and three nearby air quality monitoring sites on the 12 August 2010. As shown, the rise of $\text{SO}_2$ was simultaneously observed at three other air quality monitoring sites.
near the urban basin, suggesting regional-scale SO$_2$ impact occurred at 15:00 LT.

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